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**Mapping Exposed Silicate Rock Types and Exposed Ferric
and Ferrous Compounds from a Space Platform**

Quarterly Report for Period 8 March - 8 June 1974

EREP Investigation 444M
NASA Contract NAS9-13317

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Quarterly Report for Period 8 March - 8 June 1974

The following report serves as the fifth quarterly report for this contract, which is entitled "Mapping Exposed Silicate Rock Types and Exposed Ferric and Ferrous Compounds from a Space Platform." The financial reports have been submitted monthly under separate cover.

During this quarter, screening film was received for SL-4 data, start/stop times of S-192 data were defined for the Southern California test site, a paper was presented at the Ninth Remote Sensing of Environment Symposium which theoretically (based on laboratory spectra) predicts the usefulness of S-192 data for geological remote sensing, and ratio processing of actual S-192 line-straightened data (received from NASA for another ERIM contract) was begun for algorithm testing purposes on an area near White Sands, New Mexico. The screening film of S-192 data over Southern California of bands 6, 9, and 13-1 are encouraging from the standpoint of apparently good signal-to-noise. The Pisgah Crater test site was only partially covered by the data collected, but the remainder of the image strip provides enough geological targets to make up for the uncovered portions of the test area. The S-192 start and stop times requested for line straightening are as follows:

START 026:19:42:21 GMT

STOP 026:19:43:25 GMT

The data were collected on 26 January 1974, but the regions of interest are free of snow and clouds.

The contrast in thermal image (from band 13-1) may be good enough to enable detection of very warm geothermal areas. An ozalid color composite was made of the three channels of data to search for areas which have high visible-reflective IR albedo, but yet are warmer than average for the scene. For such anomalous areas, the above-average thermal brightness could be caused by bright-faced slopes toward the solar direction (sun elevation was approximately 45°, since data collection was at mid-day on 26 January), relatively low thermal

inertia materials, or geothermal heat sources. Geothermal sources close enough to the surface to have temperatures elevated enough to appear warm on SKYLAB S-192 images may also be accompanied by relatively bright clays and possibly iron oxides, both of which can be products of hydrothermal alteration. Thus far, two anomalies have been found in the color composite which have high visible-reflective IR albedo and yet are warmer than average. [Note: This is neither a sufficient nor necessary condition for geothermal sources; it only eliminates the low albedo-high temperature targets.] One area is small (about 1 km diameter) and occurs about two miles north of a concealed portion of the Pinto Mountain fault, just west of Twenty-Nine Palms, California. From a recent geologic map, there is no evidence of either an appreciable slope or geothermal activity. The other area, for which a geologic map has been ordered, is in the Southern end of Palen Valley. The 1:250,000 topographic map for the Salton Sea region shows that sand dunes and Palen Dry Lake are in the general vicinity of this anomaly, which would suggest that a sunward-facing slope of bright material (such as sand) may be causing the anomaly. These areas will be checked in the field within the next six months, if possible. In addition there are warm places along the San Andreas fault, but not even low albedo could be excluded from the possible causes of the slightly above-average temperatures of those places by this rather crude color composite method.

The paper presented at the Symposium (Ann Arbor, Michigan) in April, entitled "SKYLAB S-192 Ratio Codes of Soil, Mineral, and Rock Spectra for Ratio Image Selection and Interpretation," is reprinted at the end of this quarterly report. It is a limited systems study of the S-192 scanner for geological remote sensing, using 211 laboratory spectra as "signatures" of various rocks, minerals, and soils. Included in it are the following significant items of information:

- 1) Prioritization of the single channels of SKYLAB for compositional mapping.
- 2) Selection of twelve (of a possible sixty-six) best ratios for compositional mapping.
- 3) Ratio codes of the top twelve spectral ratios for all 211 samples, for choosing the best ratio image to enhance particular geologic targets.
- 4) When to use single channel inputs and when to use ratio inputs for automatic recognition.
- 5) What geologic classes SKYLAB S-192 data are expected to separate well and poorly.

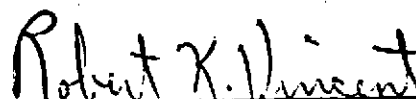
This paper should be helpful to all investigators using ratioing for compositional mapping, and may be of assistance in the selection of spectral channels for future satellites. It is the most significant work reported thus far in this contract.

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Finally, it has been decided to use line-straightened S-192 data already supplied to ERIM under another contract (NAS9-13386) for the purpose of testing the ratio algorithms to be used later with Southern California data of SL-4. The in-house data (White Sands, New Mexico) are rather bland in geological content compared with the Southern California test site, but it will be used to prepare the way for the SL-4 data, whenever it becomes available. More will be reported on this subject next quarter.

Respectfully submitted:



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attachment

SKYLAB S-192 RATIO CODES OF SOIL, MINERAL, AND
ROCK SPECTRA FOR RATIO IMAGE SELECTION AND INTERPRETATION

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ABSTRACT

The SKYLAB S-192 multispectral scanner has twelve channels in the 0.4 - 2.5 μm wavelength region and a single broad-band thermal infrared channel, intended solely for temperature measurements. Since spectral ratio images have been shown with other scanners to be useful for enhancing certain geologic targets and for suppressing unwanted atmospheric and solar illuminations across the scanned scene, ratioing will be employed for similar purposes with SKYLAB data. There are 66 spectral ratios (excluding reciprocals) that can be constructed from the twelve visible-reflective infrared channels of the S-192 scanner. Clearly, it is not economically feasible to produce all possible ratio images for large blocks of SKYLAB data. To aid investigators in selecting optimum ratio images, all possible SKYLAB single channel reflectances and ratios were calculated for approximately 211 laboratory spectra of soils, minerals, and rocks. Linear discriminant analysis was used to separately rank the twelve SKYLAB channels and sixty-six ratios according to their ability to classify these laboratory samples into seventeen assigned classes. The twelve highest ranked ratios were converted to ratio codes, which compress each spectrum in the 0.4 - 2.5 μm wavelength region to a single number. The ratio codes can be used to select the ratio images which will best enhance a target of interest. They can also be easily searched for false alarm candidates (or "look-alikes") of a given target.

For automatic geologic mapping, twelve single channel inputs to linear discriminate analysis were found superior for seven geologic classes, twelve ratio inputs to the same statistical analysis were superior for six classes, and the two cases produced nearly equivalent results for four classes. However, the greater independence of ratios on environmental and instrumental noise tends to favor ratios for the latter four classes. Because of their simplicity and photo-interpretability, ratio images are recommended over automatic mapping methods for those materials which the ratio codes indicate can be enhanced by ratioing.

The results of this study indicate that the SKYLAB S-192 scanner should be most useful for mapping products of hydrothermal alteration, weathering, and evaporation, but that discrimination among igneous rocks will be difficult. This suggests that S-192 data should be quite useful for mineral exploration, but the full potential of general surface geologic mapping by satellite will probably not be reached until high spatial resolution ($\leq 30\text{m}$) scanners with multiple-channel thermal infrared capability have been orbited.

INTRODUCTION

The first ratio image from multispectral scanner data was produced in 1970 from aircraft data and reported in the Seventh International Symposium on Remote Sensing of Environment [1]. In that experiment, the radiances of two thermal infrared channels were ratioed, resulting in a ratio image which was crudely correlated with SiO_2 content of exposed rocks. Since then, spectral ratio imaging has been employed by several investigators of various disciplines. Some of the most recent geological applications of ratio imaging have been for the purpose of enhancing iron oxides in ERTS MSS data [2,3]. Plans are now being made by many geological investigators to utilize the SKYLAB S-192 multispectral scanner for geological remote sensing. They will be attempting to enhance particular minerals and rocks and to perform general geological mapping. Some of this processing will involve spectral ratioing. A typical number of single channels or ratios utilized for producing automatic recognition maps is six. However, there are twelve S-192 channels in the 0.4 - 2.5 μm spectral region, from which sixty-six unique (non-reciprocal) spectral ratios can be formed. There is a clear need to prioritize both the single channels and ratios according to their ability to discriminate among rocks, minerals, and soils for geological remote sensing experiments from SKYLAB S-192 data.

The purpose of this paper is to provide qualified answers to the following questions concerning the twelve visible-reflective IR channels of the SKYLAB S-192 scanner, on the basis of 211 laboratory spectra of rocks, minerals, soils, and some vegetation:

- 1) What are the best single channels with which to produce automatic recognition maps for geological remote sensing?
- 2) What are the best spectral ratios with which to produce automatic recognition maps for geological remote sensing?
- 3) What materials can be enhanced in a single spectral ratio image, and which ratio should be chosen for each of these materials?
- 4) For which classes of geological targets do single channels provide better multispectral discrimination than spectral ratios, and vice versa?
- 5) What geologic classes will SKYLAB S-192 be most capable of discriminating?

The qualifications to the answers to these questions will be threefold: no environmental (atmosphere, solar illumination, etc.) or instrumental (electrical noise, gain factors, etc.) effects are considered; the 211 laboratory spectra utilized represent available data and do not provide good statistical samples of all target classes; and the seventeen target classes defined for this study are somewhat subjective.

METHODOLOGY

The Earth Resources Spectral Information System (ERSIS), created for NASA by ERIM, contains over 3,000 reflectance spectra of natural materials [4,5,6]. ERSIS was searched for all spectra of rocks, minerals, and soils that spanned the 0.4 - 2.5 μm wavelength region and which represented the more realistic grain-sized samples (samples with particle diameters exclusively less than 74 μm were omitted). To this, three vegetation spectra (coarse grass, clover, and dead grass) were added. The resulting 211 spectra, which hereafter will be called the geological data collection, were divided into seventeen classes, as shown by Table 1. The mineral classes, 1-9 and 17, represent spectra of pure minerals.

Linear discriminant analysis [7] was the method selected for prioritizing the S-192 channels and spectral ratios according to their ability to discriminate among the seventeen classes of Table 1. This method, which assumes equal covariances, is not expected to be as powerful as the STEP-L program [8] of ERIM, which chooses best channels by minimizing the average pairwise probability of misclassification, but the linear discriminant analysis method is currently more amenable to the use of laboratory spectra as signatures. Given the known reflectance spectra of seventeen classes of materials, an unidentified reflectance spectrum can be classified as one of these classes by calculating its discriminant function for each of the seventeen classes, and assigning the unknown spectrum to the class for which it had the largest discriminant function. The equation for the discriminant function D_{ik} of the i th target class for the k th sample spectrum in the i th target class is given by

$$D_{ik} = a_{i0} + \sum_{j=1}^n a_{ij} C_{jik} \quad (1)$$

where

C_{jik} = spectral parameter of the k th sample spectrum of the i th class for the j th channel (or j th ratio)

n = number of channels (or ratios)

a_{i0} = constant for the i th class

a_{ij} = discriminant coefficient for the i th class and j th channel (or ratio)

When only single channels are used as inputs to this classification scheme, C_{jik} is the average laboratory-measured reflectance of the i , k th sample in the spectral region covered by the j th multispectral scanner channel. When spectral ratios are the only inputs, C_{jik} is the quotient resulting from a division of laboratory-measured reflectances in the two wavelength regions covered by the two multispectral scanner channels used to form the j th ratio. There is only one set of coefficients (a_{i0} and a_{ij} 's) for each target class. When an unknown geological sample is to be classified, its

TABLE 1
ROCK, MINERAL, SOIL, AND VEGETATION CLASSES

| CLASS NUMBER (1) | NAME | NUMBER IN CLASS (K1) |
|------------------------|--|----------------------------|
| 1 | Oxides and hydroxides (Excluding Common Iron Oxides) | 27 |
| 2 | Sulfur, Sulfides, and Sulfates | 24 |
| 3 | Halides | 6 |
| 4 | Phosphates | 5 |
| 5 | Carbonates | 29 |
| 6 | Clay Minerals | 8 |
| 7 | Quartz and Feldspars | 19 |
| 8 | Ferromagnesian Minerals | 35 |
| 9 | Minor Silicate Minerals | 8 |
| 10 | Felsic Rocks and Chert | 7 |
| 11 | Intermediate Rocks | 5 |
| 12 | Basic and Ultrabasic Rocks | 9 |
| 13 | Clay Soils | 3 |
| 14 | Loam Soils | 13 |
| 15 | Sandy Soils and Sand | 6 |
| 16 | Vegetation | 3 |
| 17 | Common Iron Oxides | 4 |

reflectance spectrum is measured, C_{jik} 's are calculated ($C_{jik} = C_{jk}$ for the unknown sample), and substituted into equation (1) for each target class. Hence, for 17 target classes there are seven-teen discriminant functions which are calculated. The sample is then classified as belonging to that target class for which the discriminant function is greatest. Geometrically, a discriminant function for a given class describes where the unknown sample spectrum vector lies in n -space, relative to a hyperplane that separates the volume which encloses the given class from the volumes defining each of the other classes.

Although linear discriminant analysis was originally devised as a classification method, it can also be useful for selecting the best channels or ratios for discriminating among the defined target classes (17 in this case). [Note: The method and results described here for ranking single channels and ratios are somewhat different (and better) than those given in the oral presentation of this paper. The following describes the subroutine SEPARATE in the University of Michigan's Statistical Research Laboratory's MIDAS program collection.] The best single channel (or ratio) is chosen on the basis of the largest F -statistic, where

$$F_j = \frac{\frac{1}{I-1} \sum_{i=1}^I K_i (\bar{X}_{ij} - \bar{X}_j)^2}{\frac{1}{(N-I)} \sum_{i=1}^I \sum_{k=1}^{K_i} (X_{ijk} - \bar{X}_{ij})^2} \quad (2)$$

where

F_j = F -statistic for j th single channel (or ratio)

I = no. of target classes (17)

K_i = no. of samples in i th target class

N = total no. of samples (211) in the geological data collection

X_{ijk} = observed reflectance (or ratio value) in the j th channel (or ratio) of the k th sample in the i th class

$\bar{X}_{ij} = \frac{1}{K_i} \sum_{k=1}^{K_i} X_{ijk}$ = ave. observed reflectance (or ratio value) of the j th channel (or ratio) over the i th class

$\bar{X}_j = \frac{1}{N} \sum_{i=1}^I \sum_{k=1}^{K_i} X_{ijk}$ = ave. observed reflectance (or ratio value) of the j th channel (or ratio) over all samples in the geological data collection

Once the best channel (or ratio) is chosen, a discriminant function D_{ijk} is calculated from equation 1 for each sample in the geological data collection, using the best channel (or ratio) and one of the remaining channels (or ratios) denoted by j' . These D_{ijk} are substituted for the observed reflectances (or ratios), X_{ijk} , in equation 2 and $F_{j'}$ is calculated. For all the other remaining j' channels, $F_{j'}$ are calculated in like manner. The remaining channel with the highest $F_{j'}$ is chosen as second best. This procedure is reiterated, calculating D_{ijk} 's with the best two channels and one of the other remaining channels. The best linear combination of channels (or ratios) is found through this regression procedure. After each new channel is selected, the significance of all channels are tested. If a previously chosen channel becomes less significant than a user-specified level, it is discarded and becomes a member of the "remaining" channels once again. The regression proceeds until a user-specified significance level of the final F -statistic is reached. In this paper, a new channel was selected only if it had a significance level less than 0.10, and a previously chosen channel was rejected if its significance level rose above 0.20.

RANKINGS OF S-192 SINGLE CHANNELS AND SPECTRAL RATIOS

The twelve single channels of the S-192 multispectral scanner were prioritized according to the above procedure, with the resulting rankings shown in Table 2. To do this, the average reflectances of all 211 spectra were calculated for the twelve visible-reflective IR S-192 channels, and these average reflectances were substituted into equation 1 as $C_{i,j}$. As an aid to physically understanding the meaning of these rankings, Figure 1 shows the reflectance spectra of a few minerals in the 0.4 - 2.5 μm spectral region. Spectral regions covered by SKILAB S-192 channels (channels 1-7 are not shown individually) and ERTS MSS channels are shown. The highest ranked channel (2.10 - 2.34 μm) monitors the spectral region in which reflectance minima caused by carbonate (CO_3) and hydroxyl (OH^-) ions occur. The second-ranked channel covers the 0.93 - 1.05 μm spectral region which, according to Figure 1, is primarily monitoring the absorption bands (reflectance minima) produced by transition metal ions (primarily Fe^{2+} , Fe^{3+} , and Cu^{2+}). It is not surprising that the lowest ranked channel covers the 0.42 - 0.45 μm region, because the spectra in Figure 1 show the least spectral contrast from one another in this violet region and not much new information can be obtained from this channel that is not already available from the third ranked channel. It is significant for multispectral scanners that three of the four top-ranked channels are either beyond or on the extreme edge of the spectral range of photographic film.

To find best ratios, all sixty-six non-reciprocal spectral ratios were calculated for the 211 sample spectra, and these were substituted as $C_{i,j}$ into equation 1. Resulting from the same procedure described above, the twelve highest ranked spectral ratios are shown in Table 3. Seven of the top twelve are ratios of adjacent or once-removed channels. Since ratios are a form of non-linear processing, it is not surprising that the highest ranked ratios involve some of the lower ranked single channels. It is significant that the top twelve spectral ratios utilize only ten of the twelve single channels. Channels 1 and 6 were excluded from use in the twelve highest ranked spectral ratios.

SKILAB S-192 RATIO CODES

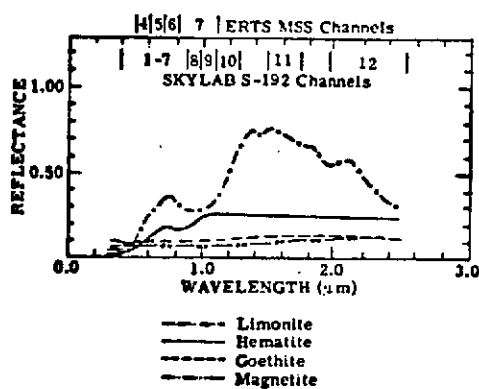
The previous section answered the first two questions posed in the Introduction. The third question will now be addressed: what materials can be enhanced in a single spectral ratio image, and which ratio should be chosen for each of these materials? Because this question requires the reader to examine spectral information from practically all of the 211 spectra in the geological data collection, a form of data compression was instituted. All twelve of the ratios shown in Table 2 were calculated for each of the 211 spectra. For each spectral ratio, a histogram of number of spectra versus ratio value was plotted. From this, ratio ranges were defined such that the spectral curve "population" was divided into deciles, and the ratio range corresponding to the first decile (the 10% of spectral curves with the lowest ratio values) was coded with "0", the second decile with "1", etc., on up to the decile with the highest ratio values, which was coded with "9". The resulting information was used to form a twelve-digit ratio code for each spectral curve, with each digit position describing a different ratio (the highest to lowest ranked ratios go from left-to-right in the ratio code). Table 4 gives the ratio ranges associated with each decile and Table 5 lists the twelve-digit ratio codes for all 211 spectral curves in the geological data collection. The ratios in Tables 4 and 5 correspond to the twelve ratios in Table 3.

Figure 2 gives an example of how to interpret ratio codes. Hematite has a "9" in the third digit position, which means that in an $R_{8,4}$ ratio image (channel 8 divided channel 4 for the S-192 scanner), hematite would appear brighter than 90% of the samples represented by the 211 spectra in the geological data collection. A "0" in the second digit position indicates that in a ratio image of $R_{3,2}$ hematite would appear darker than 90% of the other samples.

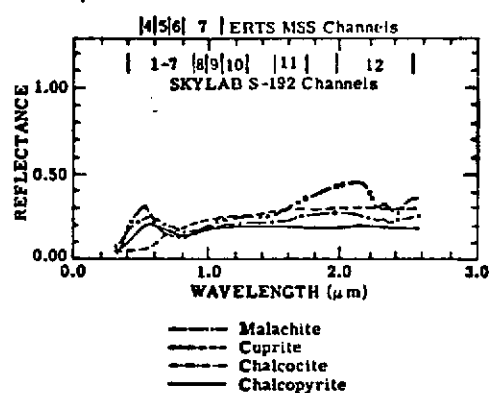
If the assumption could be made that the population of the geological data collection represents well the relative amounts of each target class present in a typical scene, then the phrase "of the other samples represented by the 211 spectra in the geological data bank" could be replaced by "of the materials in the scene". The geological data bank population described in Table 1 probably does not resemble any geological scene closely, but there appear to be enough similarities with arid, and semi-arid terrain to make a loose interpretation of this type useful. Mindful that this gross assumption has been made, for arid and semi-arid terrain it is possible to get an idea of the relative brightness of a given material in each of twelve ratio images by examining the twelve-digit ratio code of that material. The definition of enhancement for the case of ratio images is taken here to mean that a ratio image can be produced in which the target of interest is one of the brightest or darkest objects in the scene. Therefore, those ratios for which a material has a "9" or "0" ratio code are recommended for enhancing that material. In the case of hematite, for example, the $R_{8,4}$, $R_{7,5}$, $R_{8,5}$, $R_{8,7}$, and $R_{3,2}$ ratios would be recommended. Approximately 36% of the 211 samples in the geological data bank have neither a "9" nor "0" in their twelve-digit ratio codes, however. For these materials, ratio image enhancement is less useful. However, the ratio codes should be helpful in interpreting individual ratio images in general, even for those materials not enhanced in the ratio image of interest. For instance, the ratio codes can be easily searched for false alarm

FIGURE 1

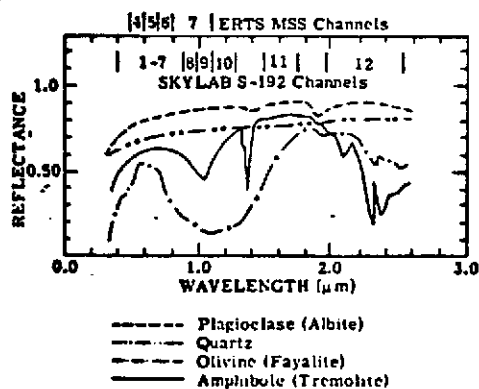
SPECTRAL REFLECTANCE OF IRON OXIDES, ORE MINERALS, AND SILICATE MINERALS
ERTS and SKYLAB Multispectral Scanner Channels are indicated on each graph



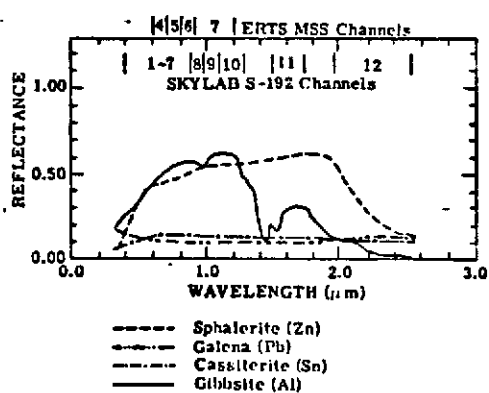
IRON OXIDES



COPPER ORE MINERALS



EXAMPLES OF SILICATE MINERALS



LEAD, ZINC, TIN, AND ALUMINUM ORE MINERALS

FIGURE 2

TWELVE-DIGIT SKYLAB S-192 RATIO CODE FOR HEMATITE

| | | | | | | |
|------------------------------------|-----------|------------|------------|-----------|-----------|-----------|
| | $R_{7,5}$ | $R_{8,4}$ | $R_{12,1}$ | $R_{4,2}$ | $R_{7,2}$ | $R_{8,5}$ |
| | ↑ | ↑ | ↑ | ↑ | ↑ | ↑ |
| Hematite (74 - 250 μ m, Minn.) | 9 0 | 9 1 | 6 8 | 5 8 | 8 8 | 9 9 |
| | ↓ | ↓ | ↓ | ↓ | ↓ | ↓ |
| | $R_{3,2}$ | $R_{10,9}$ | $R_{7,3}$ | $R_{4,3}$ | $R_{7,4}$ | $R_{8,7}$ |

where $R_{i,j} = \frac{\rho_i}{\rho_j}$

and ρ_i and ρ_j are the integrated reflectances of a laboratory spectrum of hematite in the i th and j th spectral channels of the SKYLAB S-192 multispectral scanner, respectively.

Each digit corresponds to a ratio range that defines a decile of the population in the data bank. For instance, a 9 in $R_{8,4}$ means that hematite would be brighter in an $R_{8,4}$ ratio image than 90% of the other materials in the data bank.

TABLE 2

RANKING OF SKYLAB S-192 CHANNELS FOR PRODUCING AUTOMATIC
RECOGNITION MAPS OF ROCK, MINERAL, AND SOIL CLASSES
(BASED ON DISCRIMINANT ANALYSIS OF 211 LABORATORY SPECTRA)

| RANK | CHANNEL | |
|------|---------|-----------------------|
| | NUMBER | WAVELENGTH (μ m) |
| 1 | 12 | 2.10 - 2.34 |
| 2 | 8 | 0.93 - 1.05 |
| 3 | 2 | 0.45 - 0.50 |
| 4 | 11 | 1.55 - 1.73 |
| 5 | 5 | 0.60 - 0.65 |
| 6 | 4 | 0.54 - 0.60 |
| 7 | 7 | 0.77 - 0.89 |
| 8 | 9 | 1.03 - 1.19 |
| 9 | 10 | 1.15 - 1.28 |
| 10 | 6 | 0.65 - 0.73 |
| 11 | 3 | 0.50 - 0.55 |
| 12 | 1 | 0.42 - 0.45 |

TABLE 3

TWELVE BEST RATIOS FOR PRODUCING AUTOMATIC RECOGNITION MAPS
OF ROCK, MINERAL, AND SOIL CLASSES FROM SKYLAB S-192 SCANNER DATA
(BASED ON DISCRIMINANT ANALYSIS OF 211 LABORATORY SPECTRA)

| RANK | RATIO |
|------|---|
| 1 | $R_{7,5} = \frac{L_7 (0.77 - 0.89 \mu\text{m})}{L_5 (0.60 - 0.65 \mu\text{m})}$ |
| 2 | $R_{3,2} = \frac{L_3 (0.50 - 0.55 \mu\text{m})}{L_2 (0.45 - 0.50 \mu\text{m})}$ |
| 3 | $R_{8,4} = \frac{L_8 (0.93 - 1.05 \mu\text{m})}{L_4 (0.54 - 0.60 \mu\text{m})}$ |
| 4 | $R_{10,9} = \frac{L_{10} (1.15 - 1.28 \mu\text{m})}{L_9 (1.03 - 1.19 \mu\text{m})}$ |
| 5 | $R_{12,11} = \frac{L_{12} (2.10 - 2.34 \mu\text{m})}{L_{11} (1.55 - 1.73 \mu\text{m})}$ |
| 6 | $R_{7,3} = \frac{L_7 (0.77 - 0.89 \mu\text{m})}{L_3 (0.50 - 0.55 \mu\text{m})}$ |
| 7 | $R_{4,2} = \frac{L_4 (0.54 - 0.60 \mu\text{m})}{L_2 (0.45 - 0.50 \mu\text{m})}$ |
| 8 | $R_{4,3} = \frac{L_4 (0.54 - 0.60 \mu\text{m})}{L_3 (0.50 - 0.55 \mu\text{m})}$ |
| 9 | $R_{7,2} = \frac{L_7 (0.77 - 0.89 \mu\text{m})}{L_2 (0.45 - 0.50 \mu\text{m})}$ |
| 10 | $R_{7,4} = \frac{L_7 (0.77 - 0.89 \mu\text{m})}{L_4 (0.54 - 0.60 \mu\text{m})}$ |
| 11 | $R_{8,5} = \frac{L_8 (0.93 - 1.05 \mu\text{m})}{L_5 (0.50 - 0.55 \mu\text{m})}$ |
| 12 | $R_{8,7} = \frac{L_8 (0.93 - 1.05 \mu\text{m})}{L_7 (0.77 - 0.89 \mu\text{m})}$ |

TABLE 4
RATIO RANGES FOR S-192 RATIO CODES

| | <u>CODE</u> | | | | | | | | | | |
|--------------------|-------------|-------|-------|-------|-------|-------|-------|-------|-------|-------|--------|
| RATIO | 0 | 1 | 2 | 3 | 4 | 5 | 6 | 7 | 8 | 9 | |
| R _{7,5} | 0.392 | 0.940 | 1.021 | 1.063 | 1.080 | 1.123 | 1.157 | 1.229 | 1.311 | 1.499 | 11.086 |
| R _{3,2} | 0.754 | 1.011 | 1.028 | 1.045 | 1.066 | 1.082 | 1.116 | 1.155 | 1.217 | 1.397 | 1.943 |
| R _{8,4} | 0.277 | 0.839 | 0.986 | 1.049 | 1.110 | 1.177 | 1.255 | 1.419 | 1.650 | 2.589 | 7.085 |
| R _{10,9} | 0.858 | 0.972 | 0.992 | 1.000 | 1.008 | 1.025 | 1.034 | 1.051 | 1.082 | 1.199 | 1.453 |
| R _{12,11} | 0.162 | 0.524 | 0.727 | 0.828 | 0.891 | 0.937 | 0.998 | 1.040 | 1.105 | 1.267 | 1.955 |
| R _{7,3} | 0.336 | 0.876 | 1.036 | 1.095 | 1.166 | 1.257 | 1.381 | 1.528 | 1.864 | 2.808 | 8.485 |
| R _{4,2} | 0.743 | 1.035 | 1.065 | 1.090 | 1.116 | 1.163 | 1.214 | 1.310 | 1.515 | 1.860 | 3.742 |
| R _{4,3} | 0.893 | 1.003 | 1.027 | 1.036 | 1.050 | 1.067 | 1.095 | 1.147 | 1.192 | 1.365 | 2.507 |
| R _{7,2} | 0.382 | 0.949 | 1.093 | 1.139 | 1.227 | 1.350 | 1.512 | 1.768 | 2.330 | 3.521 | 15.684 |
| R _{7,4} | 0.319 | 0.854 | 1.006 | 1.063 | 1.106 | 1.162 | 1.232 | 1.360 | 1.545 | 2.124 | 6.215 |
| R _{8,5} | 0.331 | 0.884 | 0.990 | 1.051 | 1.090 | 1.149 | 1.191 | 1.301 | 1.406 | 1.672 | 12.638 |
| R _{8,7} | 0.589 | 0.922 | 0.968 | 0.991 | 1.003 | 1.023 | 1.038 | 1.067 | 1.105 | 1.181 | 1.840 |

TABLE 5

S-192 RATIO CODES OF GEOLOGICAL DATA COLLECTION

| ERSIS DOCUMENT NO. | SAMPLE DESCRIPTION | RATIO CODE |
|--------------------|----------------------------|--------------------------|
| B30000001 | ATACAMITE, NATURAL SURFACE | 281191 734135 |
| B30000002 | SYLVITE 250-1200 MICR | 424005 454432 |
| B30000003 | SYLVITE 74-250 MICR | <u>212004</u> 133410 |
| B30000004 | ANNABERGITE 74-250 MICR | 445021 001268 |
| B30000005 | ANNABERGITE 250-1200 MICR | 153021 561159 |
| B30000006 | CARNOTITE | 797127 978677 |
| B30000007 | VIVIANITE 74-250 MICR | 394196 887445 |
| B30000008 | VIVIANITE 250-1200 MICR | 686697 877652 |
| B30000009 | CELEMANITE 74-250 MICR | 534004 234442 |
| B30000010 | CELEMANITE 250-1200 MICR | <u>332003</u> 343321 |
| B30000011 | ULEXITE 74-250 MICR | 212002 122222 |
| B30010001 | GYPSUM DUNE SAND | 453002 303332 |
| B30010002 | FLAYA CRUST | 887627 887787 |
| B30010003 | BASALT, MALPAIS FLOW | 747776 456788 |
| B00830001 | CLAY, CUIBBO GRAVELLY | DRY 678328 888663 |
| B00830017 | LCAM, CLARION | DRY 979878 778999 |
| B00830021 | LCAM, HERRADURA PURE SILT | DRY 697728 999751 |
| B00830025 | LCAM, AGUAY SILT | DRY <u>688258</u> 788886 |
| B00830029 | LCAM, AIKEN CLAY | DRY 898548 999872 |
| B00830033 | SAND, TEXAS DUNE | DRY 888268 888886 |
| B00830037 | LCAM, MOAULA LIGHT CLAY | DRY <u>998649</u> 999884 |
| B00830041 | LCAM, ZANESVILLE SILT | DRY 797757 988777 |
| B00830049 | LCAM, COLTS NECK TYPE | DRY 896949 999872 |
| B00830053 | LCAM, RUSTON SANDY TYPE | DRY 789826 999989 |
| B00830151 | LCAM, SAALERO HEAVY CLAY | DRY 998249 999982 |
| B00830152 | LCAM, GREENVILLE TYPE | DRY 999820 999981 |
| B00830163 | LCAM, SANTA BARBARA TYPE | DRY <u>797838</u> 999864 |
| B00830179 | LIMESTONE, SOFT AND WHITE | DRY 666456 666565 |
| B00830183 | LCAM, FLAKELY CLAY TYPE | DRY 989459 999997 |

TABLE 5 (cont.)

S-192 RATIO CODES OF GEOLOGICAL DATA COLLECTION

| ERSIS DOCUMENT NO. | SAMPLE DESCRIPTION | RATIO CODE |
|--------------------|---------------------------------|-----------------------------|
| B00830191 | LCAM, BLACK VOLCANIC SANDY DRY | <u>988548</u> 888598 |
| B00830199 | SAND, WINTHORST TYPE | DRY 988578 899899 |
| B00830203 | CLAY, ALCANTO TYPE | DRY <u>978068</u> 898897 |
| B04804001 | CALCIUM CARBONATE, GRANULAR | <u>203252</u> 011234 |
| B09000003 | AMPH., TREMOLITE 74-250 MICR | 221822 <u>232200</u> |
| B09000004 | AMPH., TREMOLITE 250-1200 MICR | 140912 443200 |
| B09000007 | AMPH., TREMOLITE 74-250 MICR | 131441 321112 |
| B09000008 | AMPH., TREMOLITE 250-1200 MICR | 131511 221100 |
| B09000011 | AMPH., ACTINOLITE 74-250 MICR | 471921 513110 |
| B09000012 | AMPH., ACTINOLITE 250-1200 MICR | 561913 514310 |
| B09000015 | AMPH., ACTINOLITE 74-250 MICR | 160920 511011 |
| B09000016 | AMPH., ACTINOLITE 250-1200 MICR | 070910 <u>621001</u> |
| B09000019 | AMPH., HORNBLende 74-250 MICR | 857945 205788 |
| B09000020 | AMPH., HORNBLende 250-1200 MICR | 123831 111149 |
| B09000023 | AMPH., HORNBLende 74-250 MICR | 141891 <u>331136</u> |
| B09000024 | AMPH., HORNBLende 250-1200 MICR | 333792 212233 |
| B09000027 | AMPH., HORNBLende 74-250 MICR | <u>626885</u> 334567 |
| B09000028 | AMPH., HORNBLende 250-1200 MICR | 142793 553224 |
| B09000031 | ANDALUSITE 74-250 MICR | 667547 777777 |
| B09000032 | ANDALUSITE 250-1200 MICR | <u>676656</u> <u>776657</u> |
| B09000035 | ANCRTHOCLASE 74-250 MICR | 767686 666678 |
| B09000036 | ANCRTHOCLASE 250-1200 MICR | 666686 666678 |
| B09000039 | BERYL 74-250 MICR | 021730 230019 |
| B09000040 | BERYL 250-1200 MICR | <u>031820</u> 220009 |
| B09000043 | PICRITE 250-1200 MICR | 374794 665247 |
| B09000045 | CHALAZITE 74-250 MICR | 545115 455555 |
| B09000048 | CILICRITE 74-250 MICR | 566725 655578 |
| B09000050 | CILICRITE 250-1200 MICR | 050860 410006 |
| B09000052 | CILICRITE 74-250 MICR | 062970 300009 |
| B09000055 | CANALRITE 74-250 MICR | 172783 765124 |

TABLE 5 (cont.)

S-192 RATIO CODES OF GEOLOGICAL DATA COLLECTION

| ERSIS DOCUMENT NO. | SAMPLE DESCRIPTION | RATIO CODE |
|--------------------|----------------------------------|---------------|
| B09000056 | DANFURITE 250-1200 MICR | 274685 775346 |
| B09000059 | CUMERTIERITE 74-250 MICR | 808418 467887 |
| B09000060 | CUMERTIERITE 250-1200 MICR | 908418 267898 |
| B09000062 | KACLINITE 74-250 MICR | 343103 454334 |
| B09000063 | KACLINITE 250-1200 MICR | 434104 454433 |
| B09000066 | TALC 74-250 MICR | 132622 332225 |
| B09000067 | TALC 250-1200 MICR | 050800 541004 |
| B09000070 | KACLINITE 74-250 MICR | 555115 555555 |
| B09000071 | KACLINITE 250-1200 MICR | 656205 555664 |
| B09000073 | MONTMORILLONITE 74-250 MICR | 786508 888761 |
| B09000074 | MONTMORILLONITE 250-1200 MICR | 796408 988751 |
| B09000077 | MONTMORILLONITE 74-250 MICR | 426728 992643 |
| B09000078 | MONTMORILLONITE 250-1200 MICR | 182716 887311 |
| B09000081 | MUSCOVITE 74-250 MICR | 535634 334456 |
| B09000089 | CLIVINE-FORSTERITE 74-250 MICR | 342533 443321 |
| B09000090 | CLIVINE-FORSTERITE 250-1200 MICR | 645516 566631 |
| B09000093 | CLIVINE-FAYALITE 74-250 MICR | 080850 750000 |
| B09000094 | CLIVINE-FAYALITE 250-1200 MICR | 080970 760000 |
| B09000102 | CRTHOCLASE 74-250 MICR | 545345 445554 |
| B09000103 | CRTHOCLASE 250-1200 MICR | 556335 555665 |
| B09000106 | ALBITE 74-250 MICR | 514364 133455 |
| B09000107 | ALBITE 250-1200 MICR | 424263 223344 |
| B09000110 | CLIGOCCLASE 74-250 MICR | 433263 333344 |
| B09000111 | CLIGOCCLASE 250-1200 MICR | 444264 444433 |
| B09000114 | ANDESINE 74-250 MICR | 566475 566507 |
| B09000115 | ANDESINE 250-1200 MICR | 475685 776457 |
| B09000118 | LAPPALEHITE 74-250 MICR | 233372 333234 |
| B09000119 | LAPPALEHITE 250-1200 MICR | 222171 211134 |
| B09000122 | BYTOWHITE 74-250 MICR | 241171 432123 |
| B09000123 | BYTOWHITE 250-1200 MICR | 261071 552111 |

TABLE 5 (cont.)

S-192 RATIO CODES OF GEOLOGICAL DATA COLLECTION

| ERSIS DOCUMENT NO. | SAMPLE DESCRIPTION | RATIO CODE |
|--------------------|-----------------------------------|---------------|
| B09000126 | PYRX., AUGITE 74-250 MICR | 061871 541114 |
| B09000127 | PYRX., AUGITE 250-1200 MICR | 050880 410013 |
| B09000130 | PYRX., DICPSIDE 74-250 MICR | 070990 510000 |
| B09000131 | PYRX., DICPSIDE 250-1200 MICR | 070990 540000 |
| B09000134 | PYRX., HEDENBERGITE 74-250 MICR | 455694 665446 |
| B09000135 | PYRX., HEDENBERGITE 250-1200 MICR | 465396 676643 |
| B09000137 | PYRX., HYPERSTHENE 74-250 MICR | 777998 777850 |
| B09000138 | PYRX., HYPERSTHENE 250-1200 MICR | 161994 665300 |
| B09000141 | PYRX., BRONZITE 74-250 MICR | 080890 870008 |
| B09000142 | PYRX., BRONZITE 250-1200 MICR | 080990 881003 |
| B09000145 | QUARTZ 250-1200 MICK | 323373 122345 |
| B09000146 | QUARTZ, MILKY 149-250 MICR | 535355 444555 |
| B09000147 | QUARTZ, MILKY 250-420 MICR | 556455 445566 |
| B09000155 | SERPENTINE 250-1200 MICK | 040620 100011 |
| B09004007 | COARSE GRASS MICR | 999119 889996 |
| B09004008 | CLOVER MICR | 999119 989998 |
| B09004009 | PINE NEEDLES MICR | 999019 999996 |
| B09004010 | LAVA, WEATHERED (FEO STAIN) | 654725 555551 |
| B09004011 | PUMICE | 787357 777776 |
| B09004013 | LAVA, UNWEATHERED | 623583 002453 |
| B09005003 | BRUCITE 74-250 MICR | 213303 243333 |
| B09005004 | BRUCITE 250-1200 MICR | 141402 443211 |
| B09005007 | CASSITERITE 250-1200 MICR | 827147 357862 |
| B09005008 | CASSITERITE 74-250 MICR | 454045 565531 |
| B09005011 | CHRYSOBERYL 74-250 MICR | 645435 444566 |
| B09005012 | CHRYSOBERYL 250-1200 MICK | 666625 656667 |
| B09005017 | CORUNDUM 74-250 MICR | 717625 005677 |
| B09005018 | CORUNDUM 250-1200 MICR | 808716 005788 |
| B09005019 | COFFITE 74-250 MICR | 909669 568999 |
| B09005020 | COFFITE 250-1200 MICK | 909669 668999 |

TABLE 5 (cont.)

S-192 RATIO CODES OF GEOLOGICAL DATA COLLECTION

| ERSIS DOCUMENT NO. | SAMPLE DESCRIPTION | RATIO CODE |
|--------------------|---------------------------|---------------|
| 809005024 | CLASPORE 74-250 MICK | 868617 657898 |
| 809005025 | CLASPORE 250-1200 MICK | 868617 667788 |
| 809005028 | GIBBSITE 74-250 MICK | 657106 666676 |
| 809005029 | GIBBSITE 250-1200 MICK | 767007 777774 |
| 809005032 | GCETHITE 74-250 MICK | 393787 958521 |
| 809005033 | GCETHITE 250-1200 MICK | 194756 887424 |
| 809005036 | FEMATITE 74-250 MICK | 909168 588899 |
| 809005039 | ILMENITE 74-250 MICK | 727576 345677 |
| 809005042 | LIMONITE 250-1200 MICK | 897919 999260 |
| 809005045 | MAGNETITE 74-250 MICK | 214294 343436 |
| 809005048 | MAGNETITE 74-250 MICK | 141291 201113 |
| 809005050 | PSILOMELANE 74-250 MICK | 404283 012355 |
| 809005051 | PSILOMELANE 250-1200 MICK | 201581 000136 |
| 809005053 | PYRCLUSITE 74-250 MICK | 456844 555479 |
| 809005054 | PYRCLUSITE 250-1200 MICK | 657836 666688 |
| 809005057 | PYRCLUSITE 74-250 MICK | 848787 667888 |
| 809005058 | PYRCLUSITE 250-1200 MICK | 819788 247899 |
| 809005060 | RUTILE 74-250 MICK | 939279 678998 |
| 809005061 | RUTILE 250-1200 MICK | 959189 779999 |
| 809005063 | RUTILE 74-250 MICK | 969859 778999 |
| 809005064 | RUTILE 250-1200 MICK | 979958 778999 |
| 809008003 | AZURITE 74-250 MICK | 103930 000079 |
| 809008004 | AZURITE 250-1200 MICK | 002930 000069 |
| 809008006 | CALCITE 74-250 MICK | 313232 112334 |
| 809008008 | CALCITE 250-1200 MICK | 313322 112345 |
| 809008010 | CALCITE 74-250 MICK | 313342 112333 |
| 809008012 | CALCITE 250-1200 MICK | 211331 111122 |
| 809008014 | CELESTITE 74-250 MICK | 212342 011223 |
| 809008016 | CELESTITE 250-1200 MICK | 202432 011222 |
| 809008018 | CELESTITE 74-250 MICK | 313433 122233 |

TABLE 5 (cont.)

S-192 RATIO CODES OF GEOLOGICAL DATA COLLECTION

| ERSIS DOCUMENT NO. | SAMPLE DESCRIPTION | RATIO CODE |
|--------------------|----------------------------|---------------|
| 809008020 | DOLomite 250-1200 MICR | 312423 122322 |
| 809008023 | MAGNESITE 250-1200 MICR | 424133 233445 |
| 809008024 | MAGNESITE 74-250 MICR | 333323 223344 |
| 809008026 | MAGNESITE 74-250 MICR | 231632 332210 |
| 809008027 | MAGNESITE 250-1200 MICR | 130712 322100 |
| 809008030 | MALACHITE 74-250 MICR | 070630 400009 |
| 809008031 | MALACHITE 250-1200 MICR | 060780 200019 |
| 809008033 | REDOCHROSITE 74-250 MICR | 404456 135710 |
| 809008034 | REDOCHROSITE 250-1200 MICR | 304537 056700 |
| 809008036 | SIDERITE 74-250 MICR | 670596 776600 |
| 809008037 | SIDERITE 250-1200 MICR | 675197 777720 |
| 809008039 | SMITHSONITE 74-250 MICR | 020930 110000 |
| 809008040 | SMITHSONITE 250-1200 MICR | 020910 000000 |
| 809008041 | STRENTIANITE 74-250 MICR | 535444 344555 |
| 809008042 | STRENTIANITE 250-1200 MICR | 535434 334555 |
| 809008044 | WITHERITE 74-250 MICR | 423253 233345 |
| 809008045 | WITHERITE 250-1200 MICR | 434444 444445 |
| 809009003 | ALUNITE 74-250 MICR | 656007 666764 |
| 809009004 | ALUNITE 250-1200 MICR | 878008 888873 |
| 809009011 | BARITE 74-250 MICR | 777677 987777 |
| 809009012 | BARITE 250-1200 MICR | 788578 888888 |
| 809009015 | CELESTITE 74-250 MICR | 707365 005776 |
| 809009016 | CELESTITE 250-1200 MICR | 808357 006897 |
| 809009027 | CINNABAR 74-250 MICR | 859559 899987 |
| 809009028 | CINNABAR 250-1200 MICR | 749559 899988 |
| 809009042 | GYPsum 74-250 MICR | 313013 122352 |
| 809009043 | GYPsum 250-1200 MICR | 222003 123521 |
| 809009046 | GYPsum 74-250 MICR | 312002 122222 |
| 809009047 | GYPsum 250-1200 MICR | 312002 112221 |
| 809009050 | GYPsum 74-250 MICR | 414013 123443 |

TABLE 5 (cont.)

S-192 RATIO CODES OF GEOLOGICAL DATA COLLECTION

| ERSIS DOCUMENT NO. | SAMPLE DESCRIPTION | RATIO CODE |
|--------------------|------------------------------|---------------|
| B09009051 | GYPSON 250-1200 MICR | 503003 013442 |
| B09009052 | GYPSON 74-250 MICR | 515014 124544 |
| B09009060 | JAROSITE | 395937 998544 |
| B09009086 | REALGAR 74-250 MICR | 889469 995587 |
| B09009087 | REALGAR 250-1200 MICR | 889579 999988 |
| B09009096 | SULPHUR 74-250 MICR | 786466 646676 |
| B09009097 | SULPHUR 250-1200 MICR | 797466 857777 |
| B09009098 | VOLCANIC SUBLIMATE | 061251 521013 |
| B09009101 | THENARDITE 74-250 MICR | 515364 124555 |
| B09009102 | THENARDITE 250-1200 MICR | 525354 224555 |
| B09012016 | HORNBLende DICRITE | 222271 121122 |
| B09012017 | GRANODICRITE | 525285 344553 |
| B09012018 | PERPHYRY DICRITE | 222142 232221 |
| B09012019 | GREY RHYCLITE | 514184 234443 |
| B09012020 | HORNBLende GRANITE | 525584 223466 |
| B09012021 | PERPHYRY ANDESITE | 334174 454432 |
| B09012024 | RYTCHWITE GABRO | 212261 001122 |
| B09012025 | PERPHYRY FELSITE | 445574 454555 |
| B09012026 | HORNBLende ANDESITE | 131262 342211 |
| B09012027 | PINK RHYCLITE | 656287 776665 |
| B09012028 | BASALT | 101171 000112 |
| B09012029 | DICRITE GRANITE | 646586 566666 |
| B09012030 | PERIDOTITE-SERPENTINITE | 101141 011111 |
| B14004082 | SAND, WHITE | 848347 336798 |
| B14004084 | LAVA | 102492 000213 |
| B14004094 | MAKLE | 302361 000222 |
| A00261001 | GREENSTONE, ALTERED BASALT | 150850 311111 |
| A00263001 | CHERT | 766536 666676 |
| A01697001 | BASALT LAVA, ALTERED SURFACE | 000261 000100 |
| A02010101 | SAND, LIGHT TAN WITH GLASS | 999440 889999 |
| A02010201 | SAND, LIGHT TAN WITH GLASS | 999755 989999 |
| A02012101 | SAND, COARSE GREY | 878567 777787 |
| A02013101 | SAND, GRAY WASHED PIT | 776746 777666 |

candidates (or "look-alikes") of a given target. Eventually, these codes may be useful for logic design and memory storage in near-real-time automatic data processors on board space shuttle or aircraft earth resources systems of the future.

PREDICTED COMPOSITIONAL REMOTE SENSING CAPABILITIES OF THE S-192 SCANNER

The final two questions to be addressed are: for which classes of geological targets are single channels better inputs than spectral ratios (and vice versa) for automatic classification by linear discriminant analysis and how well is the S-192 scanner theoretically predicted to perform for each geologic class? To answer these questions, the Mahalanobis distance, a multivariate measure of separation, was calculated for each pair of classes. The Mahalanobis distance is the square of the difference of the mean vectors of two classes weighted by the pooled variance-covariance matrix. The distance between all possible pairs of classes were calculated, along with an F-statistic and a significance level of that F-statistic which indicates how well each pair of geologic classes can be discriminated. The F-statistic here is a multivariate analog of the univariate F-statistic given in equation 2 and is proportional to the Mahalanobis distance. These statistical parameters were calculated for both single channel inputs (all twelve channels), and ratio inputs (top twelve ratios). Table 6 shows a tabulation of the significance levels for all possible pairs of geologic classes. Since this table would be symmetric about the dashed diagonal line for either single channel inputs or ratio inputs alone, the results for single channel inputs are recorded above the diagonal and for ratio inputs are recorded below the diagonal. The class numbers are the same as those in Table 1. The significance level is the probability that the same Mahalanobis distance could have been calculated from two purely random (and inseparable) geologic classes. Hence, the lower the significance level, the better the S-192 scanner should discriminate between the pair of classes. For example, Table 6 shows that for separating classes 17 (common iron oxides) and 12 (basic and ultrabasic rocks), single channel inputs would produce poor results (significance level of 0.290), whereas ratio inputs would easily discriminate between the two classes (significance level < 0.000 , indicated by a zero). Thus, ratio inputs should be employed for best automatic discrimination of these two classes by the linear discriminant analysis method. In this manner, Table 6 can answer both of the questions posed above.

In order to summarize the results of Table 6, the following evaluation system was adopted. Significance levels greater than 0.100 (which means that there is greater than a 10% chance that the same distance could have been calculated from two random classes) were interpreted to mean that the pair of classes in question could not be discriminated well. For a given class, the number of other classes which could not be discriminated well from it were counted by simply following along the proper row and column of Table 6 and counting the number of times the significance level was greater than 0.100 for single channel inputs and ratio inputs. If there were no other classes which could not be discriminated from a given class, the ability for the S-192 scanner to discriminate that class from all others was rated E for excellent. If only one or two classes could not be discriminated well from the class of interest, a rating of G for good was assigned. An M for medium was assigned for three to five classes not discriminated well, and a P for poor was assigned when S-192 had difficulty discriminating more than five other classes from the class of interest. Table 7 shows the results of this method of evaluation with each of the seventeen classes for both the case of inputting twelve single channels and the case of inputting twelve spectral ratios as C_{jik} 's in equation 1.

Using this method of evaluation, which is conservative in terms of theoretical S-192 capability (possibly nine poorly discriminated classes would have been a better upper limit for the P = poor category), there are seven geologic classes which would appear to be better discriminated by twelve single channel inputs, six classes better discriminated by twelve ratio inputs, and four classes which are discriminated approximately as well by either case. Among the six classes automatically discriminated best by ratio inputs, the common iron oxides and soils are classes for which ratios have been highly recommended by previous investigators [2,3,9] for mapping purposes. Thus, the results of Table 7 seem to be reasonable in light of past experience.

Perhaps the most prominent feature of Table 7 is the difficulty predicted for mapping igneous rocks (classes 10, 11, and 12) with the S-192 scanner. This would indeed be discouraging for geological remote sensing, were it not for the great amount of compositional information available in several medium-width spectral bands in the 8-14 μ m thermal IR region [10]. Although SKYLAB cannot collect multichannel thermal IR data (it has only one broad band thermal channel, which cannot yield significant emissivity data), the hope is that future satellite scanners will include three thermal channels, which will give assistance to geological remote sensing precisely where help is needed most. Table 7 shows, however, that the S-192 scanner should be very useful for mapping products of hydrothermal alteration, weathering, and evaporation.

Several points should be emphasized concerning the above comparisons between single channel and ratio inputs to the linear discriminant analysis equations. First, the twelve ratios utilize only ten S-192 channels of data, so in fact a ten-channel scanner is all that would be required to produce these same ratio results. Had more than twelve ratios been selected for comparison with twelve

TABLE 6
SUMMARY OF SEPARATION DISTANCES* FOR 17 GROUPS OF
ROCK, MINERAL, SOIL, AND VEGETATION CLASSES** USING 12 SINGLE CHANNEL
REFLECTANCES AND 12 BEST RATIOS OF SINGLE CHANNELS

| CLASS NUMBER | <u>1</u> | <u>2</u> | <u>3</u> | <u>4</u> | <u>5</u> | <u>6</u> | <u>7</u> | <u>8</u> | <u>9</u> | <u>10</u> | <u>11</u> | <u>12</u> | <u>13</u> | <u>14</u> | <u>15</u> | <u>16</u> | <u>17</u> |
|-----------------|----------|----------|----------|----------|----------|----------|----------|----------|----------|-----------|-----------|-----------|-----------|-----------|-----------|-----------|-----------|
| 1 | - | 0 | 0 | .005 | 0 | 0 | 0 | 0 | 0 | .055 | .290 | .059 | .091 | 0 | .512 | 0 | .232 |
| 2 | 0 | - | 0 | 0 | 0 | 0 | 0 | 0 | 0 | .001 | 0 | 0 | .011 | 0 | 0 | 0 | 0 |
| 3 | .004 | .472 | - | 0 | 0 | 0 | 0 | 0 | 0 | 0 | .001 | 0 | .004 | 0 | 0 | 0 | 0 |
| 4 | 0 | .007 | .074 | - | 0 | .001 | 0 | 0 | .001 | .047 | .120 | .023 | .138 | .001 | .111 | 0 | .013 |
| 5 | 0 | 0 | .004 | 0 | - | 0 | .002 | 0 | 0 | .006 | .002 | 0 | 0 | 0 | 0 | 0 | 0 |
| 6 | 0 | .038 | .276 | .007 | 0 | - | 0 | 0 | 0 | 0 | 0 | 0 | .069 | .002 | .001 | 0 | .003 |
| 7 | .001 | .002 | .058 | .010 | .090 | 0 | - | 0 | 0 | .417 | .006 | 0 | .006 | 0 | .003 | 0 | 0 |
| 8 | 0 | 0 | 0 | 0 | 0 | 0 | 0 | - | 0 | .012 | .034 | 0 | .001 | 0 | 0 | 0 | .006 |
| 9 | .010 | .016 | .014 | .004 | .001 | .007 | .073 | 0 | - | .016 | .005 | 0 | .013 | 0 | .080 | 0 | .012 |
| 10 | .099 | .149 | .081 | .041 | .215 | .005 | .997 | .002 | .575 | - | .870 | .157 | .476 | .025 | .770 | 0 | .081 |
| 11 | .037 | .091 | .189 | .018 | .868 | .001 | .998 | .034 | .159 | .994 | - | .997 | .287 | .004 | .304 | 0 | .322 |
| 12 | 0 | 0 | .003 | 0 | .608 | 0 | .459 | .002 | .001 | .326 | .963 | - | .059 | 0 | .050 | 0 | .290 |
| 13 | .005 | .114 | .018 | .008 | 0 | .354 | .006 | 0 | .160 | .198 | .036 | 0 | - | .825 | .706 | 0 | .176 |
| 14 | 0 | 0 | 0 | 0 | 0 | 0 | 0 | 0 | 0 | 0 | 0 | 0 | 0 | - | .522 | 0 | .119 |
| 15 | 0 | 0 | 0 | 0 | 0 | 0 | 0 | 0 | 0 | .002 | 0 | 0 | .018 | 0 | - | 0 | .418 |
| 16 | 0 | 0 | 0 | 0 | 0 | 0 | 0 | 0 | 0 | 0 | 0 | 0 | 0 | 0 | 0 | - | 0 |
| 17 | 0 | 0 | 0 | 0 | 0 | 0 | 0 | 0 | 0 | 0 | 0 | 0 | 0 | .145 | 0 | 0 | - |

*Distance is the Mahalanobis multivariate measure of separation, $V_1 - V_j$, where V_1 and V_j are the mean vectors of classes 1 and j.

**Values tabulated are the significance levels (lower significance level means higher probability of separation) of the observed F statistic; upper half for single channel separations, lower half for ratio separations.

TABLE 7

THEORETICAL COMPARISON OF SINGLE CHANNEL AND RATIO DISCRIMINATION FOR
ROCKS, MINERALS, AND SOILS FROM THE SKYLAB S-192 MULTISPECTRAL
SCANNER, EXCLUDING ENVIRONMENTAL VARIABILITY

| <u>SYMBOL</u> | <u>NO. OF CLASSES WITH SIGNIFICANCE LEVEL >0.100</u> |
|---------------|---|
| E = EXCELLENT | 0 |
| G = GOOD | 1-2 |
| M = MEDIUM | 3-5 |
| P = POOR | >5 |

| CLASS NAME | CLASS NUMBER | SINGLE CHANNEL DISCRIMINATION | RATIO DISCRIMINATION |
|--|--------------|-------------------------------|----------------------|
| OXIDES & HYDROXIDES (EXCLUDING COMMON IRON OXIDES) | 1 | M | E |
| SULFUR, SULFIDES, & SULFATES | 2 | E | M |
| HALIDES | 3 | E | M |
| PHOSPHATES | 4 | M | E |
| CARBONATES | 5 | E | M |
| CLAY MINERALS | 6 | E | G |
| QUARTZ & FELDSPARS | 7 | G | M |
| FERROMAGNESIAN MINERALS | 8 | E | E |
| MINOR SILICATE MINERALS | 9 | E | M |
| FELSIC ROCKS & CHERT | 10 | M | P |
| INTERMEDIATE ROCKS | 11 | P | P |
| BASIC & ULTRABASIC ROCKS | 12 | M | M |
| CLAY SOILS | 13 | P | M |
| LOAM SOILS | 14 | M | G |
| SANDY SOILS & SAND | 15 | P | E |
| VEGETATION | 16 | E | E |
| COMMON IRON OXIDES | 17 | P | G |

single channel results, the ratio case probably would have been improved. Second, the effects of environmental, instrumental, and bidirectional reflectance variations have been ignored, although it has been shown that spectral ratios are less affected than single channel radiances by such noise (11). For this reason, ratios are recommended over single channel inputs for the classes indicated in Table 7 as equally discriminated by both cases. Third, Table 7 shows general results for automatic discrimination via the linear discriminant analysis method (these should also be approximately true for discrimination by the maximum likelihood method). However, the capabilities of individual ratios to enhance certain class members are best assessed from the ratio codes of Table 5. The individual ratio enhancement method involves much simpler and cheaper data processing and provides a continuous gray-toned ratio image that lends itself to photogeologic techniques.

CONCLUSIONS

This has been a limited theoretical systems study of the SKYLAB S-192 multispectral scanner capabilities for geological remote sensing. Laboratory spectra of rocks, minerals, soils, and some vegetation were the sole basis for the study, which excluded environmental and instrumental sources of noise from consideration. The twelve S-192 channels between 0.4 μm and 2.5 μm were ranked according to their discrimination ability among seventeen assigned classes of geological targets. The best twelve of a possible sixty-six non-reciprocal spectral ratios were likewise selected, and a comparison was made between the use of twelve single channels and twelve ratios for automatic classification by the linear discriminant analysis method. Although the results of this study show that there were more classes that could be better discriminated by using twelve single channels than by using twelve top-ranked ratios for automatic classification, those ratios utilized only ten of the available twelve channels of information. Also, the inclusion of real-world environmental and instrumental variations would tend to favor ratios over the single channel as inputs to automatic classification methods. In general, it appears that SKYLAB will be most useful for discriminating minerals deposited by hydrothermal alteration, weathering, and evaporation. Discrimination among igneous silicates appears to be difficult with the S-192 scanner, but future scanners can probably overcome these difficulties with multiple channels in the 8 - 14 μm thermal infrared spectral region.

Because of their relatively simple production and photointerpretability, ratio images are recommended over automatic discrimination methods for those geological targets (indicated by the ratio codes) that can be enhanced by ratio imaging. Ratio codes were created to assist in the selection of ratio images for enhancing individual geologic targets. They can also be easily searched for false alarm candidates (or "look-alikes") of a given target. Eventually, ratio codes may be useful for logic design and memory storage in near-real-time automatic data processors on board space shuttle or aircraft earth resources systems of the future.

The results of this study could no doubt be significantly improved by the addition of considerably more spectral curves of field samples to the geological data collection.

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